Review on deep convection and thunderstorms over land and ocean

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PART I CONVECTIVE ACTIVITY IN TROPICAL AND SUBTROPICAL AREA



1.1. Introduction

Deep cumulonimbus clouds are essential to the general circulation and to the global energy balance.



Introduction: Some views on "hot towers"

> Riehl and Malkus (1958.) – "hot tower" hypothesis:

- Undiluted pseudoadiabatic updraft cores 2-4 km in diameter
- > Observation over tropical Atlantic from 1974. till now are not recorded undiluted cores:
 - Small diameter
 - Rarely observe adiabatic liquid water content
 - Smaller vertical velocities than if undilute (GARP, GATE, EMEX)



- Tropical Ocean Global Atmosphere, the Coupled Ocean Atmosphere Research Experiment (TOGA COARE): 09.02.1993.
 - Employed a ship array and intensive airborne quad-Doppler radar observations
 - Bimodal convective updraft profiles: minimum vertical velocity around 3-5 km, two maxima above and below
 - Cloud top heights often of about 15 km

- MCSs typically organized into convective bands aligned parallel (moving more slowly) or normal (moving more rapidly) to the environmental shear (like squall lines)
- The best-documented system during all of TOGA COARE used for model simulations and comparation with model results

- Model simulations:
 - Straka Atmospheric Model (SAM)
 - Modeling studies from Fierro et al. (2008 (F08), 2009 (F09), 2011 (F11)) show a bimodal convective updraft profiles
 - F08 and F09 model:

single moment 10–ICE microphysics sheme with 12 discrete bulk hydrometeor categories 750 m horizontal resolution

• F11 model:

single moment 3–ICE microphysics sheme with 5 discrete bulk hydrometeor categories

300 m horizontal resolution



> A parcel's inital height influences its maximum altitude

Lowest-origin trajectories reaching the highest peak altidudes



dw/dt = BA + PGA + RES

- Below 1 km the only positive acceleration term is PGA (pressure gradient acceleration) – larger that negative BA(buoyancy acceleration) – lower W maximum
- The dominant positive acceleration from about 1 km is BA
- Between 1-7 km PGA is negative, even larger magnitude than BA (5.5-7 km)- W minimum
- Net negative acceleration in the layer between 1.5-3.5kmis mainly duetorain loading
- Above 5-7km BA is larger than PGA-resulting in secondary updraft maximum



- De increase is associated with heating from enthalpy of freezing (between 5.5-6 km with heating rates near 300K/h) and deposition (between 7.5-10 km with heating rates near 20 K/h)
- W maximum above 8 km is due to release of latent heat due to enthalpy of freezing and deposition
- Near 7.5 km noticeable cooling from sublimation (-15K/h; slightly negative W)
- The parcel resumes rising above that level- the heating rate of deposition is able to exceed that of the cooling of sublimation

Introduction: differences between continental and oceanic convection

- Slightly smaller entrainment rates for clouds over land:
 - Increase updraft strength
 - The time for the updraft to reach 500 hPa decreases
 - The fraction falling out decreases
- Cloud condensation nuclei concentration is larger over land:
 - Supercooled water content increase resulting in more enthalpy of freezing



Differences between continental and oceanic covection



- 1.2. Distribution of the most intense thunderstorms in tropical and subtropical area
 - Focus is on the most extreme events using several proxies for convective intensity
 - Purpose: determine quantitative frequency and intensity of storms
 - Seasonal and diurnal cycle of the most extreme categories



Data and methods

- The Tropical Rainfall Measuring Mission (TRMM) satellite – launched in late 1997.
 - Domain: 36°N 36°S

Instruments:

- Precipitation radar (PR)
- Passive microwave imager (TMI)
- Visible and infra red scanner (VIRS)
- Lightning imaging sensor (LIS)



- Proxies for intensity:
 - The higher the height attained by the 40-dBZ level in a PF (precipitation feature), the more intense the storm
 - The lower the minimum brightness temperature atained in a PF at 37 and 85 GHz, the more intense the storm
 - The greater the lightning flash rate attained in PF, the more intense the storm
- Database: all TRMM PFs greater then four pixels in size (>75 km²) from 01.01.1998. to 31.12.2004.

Locations of intense convective events



Seasonal cycle of the two most extreme categories



Diurnal cycle of the three most extreme categories

strong afternoon 14
maximum over land
peaks at same time but 12
of rainfall over land

broad nocturnal maximum over oceans
amplitude tends to be somore of a broad peak throughout the night, while that of rainfall has its peak near sunrise

 no obvious explanation for the stronger peak in the radar proxy



SWAP 2011.

Extreme events observed with 85 GHz passive microwave channels

 On the Special Sensor Microwave Imager (SSM/I on the F-14 satellite)

- Sun-synchronous polar orbit
- Ascending nodes near 1900 local time (LC)
- Descending nodes near 0700 LC
- Polarization-corrected temperature (PCT):

$PCT = 1.818T_{Bv} - 0.818 T_{Bh}$

TBh-horizontally polarized brightness temperature

T_{Bv}-vertically polarized brightness temperature

 thresholds are different than that of TRMM database because the SSM/I's reduced spatial resolution

Extreme events at 85 GHz observed globally from 2004

- North American maximum extends into Canada
- some intense storms observed across Eurasia

 similarity of the distributions – TRMM PR domain does capture the global distribution of extreme convective storms

at sunset, 60% more MCSs occured over land, both tropical and subtropical
at sunrise, 25% more MCSs occurred over the subtropical oceans, and 40% more MCSs occurred over the tropical oceans





SWAP 2011.

Part II CONVECTIVE ACTIVITY IN MIDDLE LATIDUDES



2.1. Characteristics of European mesoscale convective systems

Database:

- Meteosat infrared images with a 30-min time resolution and a 6 km x 6 km spatial resolution
- Five warm seasons (from April to September, 1993-1997)

MCS:

- Temperature treshold=-45°C
- Area threshold=1000 km²
- At some stage area threshold = 10 000 km²



Geographical domain



Density map of MCSs triggering



Density map of MCS occurrences



Monthly density maps of MCS occurrence

- April only few MCSs over Europe
- From May to August number of MCSs is remarkably constant
- Global south-westward shifting of MCS activity
- In August, many MCSs trigger over the Pyrenees and Spain
- In September, MCS occurrences are mainly located over the Mediterranean Sea and north Africa







MCS with short durations tend to trigger later in the afternoon then MCS with long durations Afternoon peak is less pronounced for long-living MCSs

Triggering time of MCSs with very long durations is less sensitive to the diurnal heating

3. Conclusion

PART I:

- A tropical Hot Towers should be redefined as any (not undilute) deep convective cloud with a base in the boundary layer and reaching near the upper-tropospheric outflow layer
- Bimodal convective updraft profiles:

- two W maximum (below 3 km and above 5-7 km), W minimum (above 3 km)
- Decrease in Oe above cloud base results from entrainment of lower ambient Oe air
- Increase in Oe above 6-8 km was a consequence of freezing and deposition
- Slightly smaller entrainment rates for clouds over land than those over the ocean have a very large effect on updrafts (large differences in convective intensity, ice scattering and flash rates)

- Continental thunderstorms tended to be more intense than oceanic
- The strongest convective storms are often found in semiarid regions
- Heavy rains of the oceanic ITCZ, western Amazonia, southeast Asia and Indonesia have relatively few severe storms
- Storms are more frequent over the oceans at sunrise and over the continents at sunset



PART II:

- MCSs are mainly continental (some MCS triggering is observed during second half of August and September over the western Mediterranean Sea)
- Local maxima of MCS triggering are observed near all mountain ranges – strongly correlated with orography
- On average, a theoretical 'typical European MCS' moves to east-north-east, triggers near 3 p.m. LST, lasts around 5.5 hours and dissipates near 9 p.m. LST
- The diurnal cycle is in phase with the diurnal radiative heating (except for around 20% MCSs)



Thank you for your attention!

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